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Yamakawa A, Nagano K, Ukachi M, Onishi K, Yamashita K, Shibata T, Takamiya K, Kani T, Bérail S, Donard OFX and Amouroux D (2021) Sr Isotopic Composition of NIES Certified Reference Material No. 28 Urban Aerosols. Front. Environ. Chem. 2:771759. doi: 10.3389/fenvc.2021.771759 An interlaboratory study of the National Institute for Environmental Studies (NIES) certified reference material (CRM) No. 28 Urban Aerosols collected from the filters of a central ventilating system in a building in the Beijing city center from 1996 to 2005 was performed to obtain an information value of the Sr isotopic composition. The Sr isotopic composition was measured using multi-collector-inductively coupled plasma-mass spectrometry (MC-ICP-MS) to confirm the CRM's within- and between-bottle homogeneity, and the results showed a ⁸⁷Sr/⁸⁶Sr ratio of 0.710227 ± 0.000019 (2SD, n = 18). The Sr isotopic compositions were intercompared using thermal ionization mass spectrometry (TIMS), which showed good agreement with values obtained at NIES. Subsequently, a consistent ⁸⁷Sr/⁸⁶Sr ratio was observed between two dissolution (hotplate *vs.* high-pressure bomb) and Sr separation (Sr spec resin *vs.* cation exchange resin) methods. To validate and reproduce the accuracy of our analytical methods, the Sr isotopic compositions of secondary reference materials, JB-1b and JA-2, were also measured. Our results showed that NIES CRM No. 28 is appropriate for the quality control of Sr isotopic measurements of particulate matter analyses for environmental and geochemical studies.

Keywords: Sr isotopes, atmospheric particles, MC-ICP-MS, TIMS, NIES CRM

INTRODUCTION

Atmospheric particulate matter (PM) is a complex mixture of particles with diverse chemical compositions and sizes. The chemical compositions vary depending on their source (natural *vs.* anthropogenic), environmental condition (e.g., temperature, humidity, and redox condition), and atmospheric processing (e.g., radiation, convection, and transport). PM emissions from urban and industrial areas are a critical environmental problem that affects the climate, human health, visibility, biogeochemical cycles, and atmospheric chemistry. Identifying the source(s) of emitted PM is critical for providing scientific strategies to improve air quality.

Recent studies have assessed the utility of strontium (Sr) isotopes ⁸⁷Sr/⁸⁶Sr to identify sources of atmospheric PM (e.g., Capo et al., 1998; Kanayama et al., 2002; Grousset and Biscaye, 2005; Lahd Geagea et al., 2008; Widory et al., 2010; Duarte et al., 2017). Sr has four natural isotopes: ⁸⁴Sr, ⁸⁶Sr,

⁸⁷Sr, and ⁸⁸Sr. ⁸⁷Sr is a radiogenic isotope of ⁸⁷Rb by β-decay (half-life = 4.88×10^{10} years; De Laeter et al., 2003). Due to an initial difference in Rb/Sr and age-integrated effects, the ⁸⁷Sr/⁸⁶Sr ratios of Earth surface materials vary widely (e.g., more radiogenic ancient crustal rock or less radiogenic carbonate as the parent materials of soil). Therefore, Sr isotope ratios measured at a receptor site provide clues to the source of the Sr or the mixing ratio of multiple sources. For example, a previous study on atmospheric PM in Beijing indicated that atmospheric Sr was mainly controlled by coal combustion and to a lesser extent by cement plants and/or smelters (Widory et al., 2010).

To characterize emission sources, accurate methods for determining Sr isotopic ratios in PM are required, and matrix matching between samples and standards is important for the quality control of the analysis. Currently, the NIES and other research institutes (e.g., National Institute of Standards and Technology and the European Commission Joint Research Center-Institute for Reference Materials and Measurements) provide commercially available standard reference materials; aerosol however, Sr isotopic compositions of PM reference materials have not been reported. To overcome this limitation, we aim to obtain an information value of the ⁸⁷Sr/⁸⁶Sr ratio for NIES CRM No. 28 Urban Aerosols. The objectives of this study are: (i) to determine the Sr isotope distribution within and between the bottles of the CRM using multi-collector-inductively coupled plasma-mass spectrometry (MC-ICP-MS); (ii) to confirm the consistency of the interlaboratory CRM isotopic ratio using two types of instruments, MC-ICP-MS vs. thermal ionization mass spectrometry (TIMS); and (iii) to confirm the consistency of the Sr isotopic composition using different digestion methods (hotplate vs. high-pressure bomb) and Sr separation methods (Sr spec resin vs. cation exchange resin).

MATERIALS AND METHODS

Reagents

All acids used in this study were Ultrapure-100 (Kanto Chemical Co., Inc., Japan), and aqueous solutions were prepared using Milli-Q water (Japan Millipore Ltd., Japan) at NIES. At Okayama University, EL grade 70% HNO₃ (Kanto Chemical Co., Inc., Japan) was twice distilled without dilution using a Teflon still. EL grade 36% HCl (Kanto Chemical Co., Inc., Japan) was diluted to ~6 M using Milli-Q water and then twice distilled using a Teflon still. TAMAPURE-AA-100 HClO₄ (Tama Chemicals Co., Ltd., Japan) was used without further purification.

Samples NIES CRM No. 28

NIES CRM No. 28 Urban Aerosols, collected from the filters of a central ventilating system of a building located in Beijing city center, was produced to evaluate the analytical accuracy of determining the mass fraction of selected elements (18 certified and 14 reference values) (Mori et al., 2008). The certified value of Sr is $469 \pm 16 \text{ mg/kg}$ (certificate is available on https://www.nies.go.jp/labo/crm-e/hrfba300000ble6p-att/No.

28_E.pdf). The Hg isotopic composition was also determined as an information value for the CRM (Yamakawa et al., 2020).

JA-2 and JB-1b

Sr isotopic measurements of the secondary reference, JA-2 and JB-1b, were performed using the same methods to manage the analytical accuracy of our method. These geological samples were produced at the Geological Survey of Japan in the National Institute of Advanced Industrial Science and Technology. The ⁸⁷Sr/⁸⁶Sr ratios of JA-2 and JB-1b are reported in Miyazaki and Shuto (1998) and Yuhara et al. (2000), respectively.

NIST SRM 987

NIST SRM 987 was prepared at each laboratory by the following methods. At NIES and Kumamoto University, strontium carbonate powder was dissolved in dilute HNO₃ to make a stock solution. The solution was then adjusted to 300 ng g⁻¹ in 2% HNO₃ for the isotope measurement. At Okayama University, a ~1 μ g ml⁻¹ stock solution of NIST SRM 987 was prepared by dissolving strontium carbonate powder in dilute HCl.

Sample Preparation

Sample Decomposition

Three bottles (bottle No. 044, 375, and 597) were randomly selected to assess the between- and within-bottle homogeneity of Sr isotope. The three subsamples were taken from each bottle and decomposed on a hotplate with a concentrated acid mixture of $HNO_3/HClO_4/HF$. Approximately 100 mg of powdered samples of the CRM, JA-2 and JB-1b were weighed and decomposed overnight in 5 ml of HNO_3 at 140° C. Then, 2 ml of 7 M $HClO_4$ was added and heated overnight at 200°C and dried to 1 ml. A mixture of 2 ml of 13 M HNO_3 , 1 ml of 7 M $HClO_4$, and 1 ml 30 M HF was added and heated for 2 h. The decomposed samples were heated to dryness at 200°C by step heating. The resulting sample cake was redissolved in 1 ml of 3 M HNO_3 , and the insoluble fraction was removed by centrifuging. This procedure was repeated several times, and the supernatant was subjected to the following chemical separation.

At Okayama University, the Sr isotope ratios were investigated using different acid decomposition and Sr separation methods. To obtain a representative sample, approximately 200 mg of powdered sample (bottle No. 35) was digested. The sample was dissolved in an HF/HNO₃ mixture using Teflon capsules sealed in stainless steel bombs for 72 h at 190°C. Once complete digestion was achieved, the sample was transferred to a 15 ml Teflon vial and 0.2 ml of HClO₄ was added before drying down at 120°C-200°C. The evaporated sample was further treated in 0.2 ml of HClO₄ to avoid fluoride precipitation. The sample was subsequently dissolved in HCl and ~25% split was taken for Sr isotopic analysis.

Chemical Separation

Sr was separated from the other elements, particularly Rb, in the digest using Sr spec resin (Eichrom Technologies, US) at NIES. One milliliter of Sr spec resin was packed in a size S polypropylene column (Muromachi Chemical Inc., Japan). The resin was cleaned by passing 3 ml of 3 M HCl, 18 ml of 0.05 M

Instrumentation	Nu Plasma II				
Monitored isotopes	88 (Sr), 87 (Sr, Rb), 86 (Sr, Kr), 85 (Rb),				
	84 (Sr, Kr), 83 (Kr), and 82 (Kr)				
RF power	1300 W				
Plasma gas	13.0 L min ⁻¹				
Auxiliary	0.8 L min ⁻¹				
Nebulization	1.0 L min ⁻¹				
Integration time	8 sec				
Number of cycles per block	20 cycle/block				
Number of blocks	4 blocks				
Sr concentrations of sample and	300 ng g ⁻¹				
standard					
Sensitivity of sample and standard (⁸⁸ Sr)	~10 × 10 ⁻¹¹ A				

HNO₃, and 3 ml of 3 M HNO₃. The sample dissolved in 1 ml of 3 M HNO₃ was loaded onto the column after conditioning the resin using 1 ml of 3 M HNO₃. The fraction containing Rb (and Ca, K, Mg, Ba, etc.), which began eluting immediately, was discarded by passing an additional 3 ml of 3 M HNO₃, 4 ml of 6 M HNO₃, and 1 ml of 3 M HNO₃. The Sr fraction was obtained by passing 5 ml of 0.05 M HNO₃ through a filter. During sample decomposition and chemical separation, the recovery yield was >95%. The total procedural blank for Sr was <1 ng, which was negligible compared to the sample size used in this study. For the isotopic measurement, the final solutions were diluted to a Sr concentration of 300 µg g⁻¹ using 2% HNO₃.

At Okayama University, Sr was separated by passing it through cation exchange resin in 2 M HCl (1 ml of AG50 \times 12, 200–400 mesh, packed in a size S polypropylene column (Muromachi Chemical Inc.)). To achieve complete separation of Rb, the Sr separation was repeated twice. The total procedural blank was ~30 pg, which was insignificant relative to the amount of Sr extracted.

Sr Isotope Ratio Determinations: Reproducibility and Accuracy MC-ICP-MS

The MC-ICP-MS used in this study was a Nu Plasma II (Nu Instruments, UK) at NIES. Although Sr was isolated from the matrix components and Rb by column chemistry using Sr spec resin, the signal for ⁸⁵Rb was simultaneously measured for the isobaric correction. As krypton (Kr) is present as a contaminant in the Ar plasma gas, the contribution of Kr had to be corrected. The mass numbers of 88 (Sr), 87 (Sr, Rb), 86 (Sr, Kr), 85 (Rb), 84 (Sr, Kr), 83 (Kr), and 82 (Kr) were detected using individual Faraday cups. The preamplifier gains associated with each Faraday cup were calibrated daily. The operating conditions (e.g., the torch position, Ar gas flow rates, and lens settings) were adjusted to maximize the signal intensity of ⁸⁸Sr (sensitivity of ⁸⁸Sr in 300 ng g⁻¹ was typically ~10 x 10^{-11} A). Details of the operation are summarized in Table 1. 83Kr and 85Rb were monitored for the isobaric correction of ⁸⁶Kr contribution to 86 Sr (86 Kr/ 83 Kr = 1.503), 84 Kr to 84 Sr (84 Kr/ 83 Kr = 4.955), and 87 Rb to 87 Sr (87 Rb/ 85 Rb = 0.3856). These isobaric interferences were corrected by blank subtraction. The 87 Sr/ 86 Sr ratios were corrected for mass fractionation using the exponential law relative to 86 Sr/ 88 Sr = 0.1194 (Steiger and Jäger, 1977). Because of the instability of the Ar gas flow, cone and slit degradation, and/or cup aging, Sr isotopic ratios may drift during a daylong analysis. To overcome these problems, a sample-standard bracketing technique was used. The 87 Sr/ 86 Sr values of the samples were adjusted using the NIST SRM 987 value of 0.710248 (McArthur et al., 2001).

TIMS

The Sr isotopic measurements were performed using two TIMS instruments: the TRITON (Thermo Fisher Scientific, Germany) at Kumamoto University and the Finnigan MAT 262 (Thermo Fisher Scientific) at Okayama University. Approximately 200–500 ng Sr was loaded onto a degassed W single filament along with Ta-H₃PO₄ activator, and the ⁸⁷Sr/⁸⁶Sr ratio was measured at ⁸⁸Sr intensity 3–4 × 10⁻¹¹ A. The ⁸⁷Sr/⁸⁶Sr ratios were normalized to ⁸⁶Sr/⁸⁸Sr = 0.1194 (Steiger and Jä;ger, 1977) using an exponential law. To make an interlaboratory comparison of the data possible, NIST SRM 987 was analyzed at both laboratories, and the final ⁸⁷Sr/⁸⁶Sr ratios of the samples were adjusted using the recommended NIST SRM 987 value of 0.710248 (McArthur et al., 2001).

RESULTS AND DISCUSSION

Sr Isotope Measurement Reproducibility and Accuracy

The reproducibility of the Sr isotopic compositions of the NIST SRM 987 was monitored during the study period to validate the analytical stability of our operating conditions. The ⁸⁷Sr/⁸⁶Sr ratios of 0.710243 \pm 0.000016 (2SD, n = 74), 0.710254 \pm 0.000012 (2SD, n = 15), and 0.710234 \pm 0.000022 (2SD, n =16) were obtained during the period of analysis at NIES, Kumamoto University, and Okavama University, respectively. To manage the analytical accuracy of our method, two geological standards, JA-2 and JB-1b, were analyzed at least twice on different days to monitor instrument stability at NIES. The values for the ⁸⁷Sr/⁸⁶Sr isotopic ratios of JA-2 and JB-1b were 0.706315 ± 0.000022 (2SD, n = 7) and 0.704093 ± 0.000023 (2SD, n = 8), respectively (**Table 2**). These values were identical within an acceptable error to their literature counterparts (JA-2: 0.706331 ± 0.000013 (2SD, n = 5) in Miyazaki and Shuto, 1998; JB-1b: 0.704095 \pm 0.000012 (2SD, n = 13) in Yuhara et al., 2000). The ⁸⁷Sr/⁸⁶Sr isotopic values of JA-2 and JB-1b, same aliquots prepared at NIES, showed 0.706311 ± 0.000007 and 0.706325 ± 0.000007 for JA-2, and 0.704083 ± 0.000007 and 0.704084 ± 0.000006 for JB-1b at Kumamoto University (the analytical error is described in 2SE, and 150 ratios were taken in a single measurement) (Table 2). The Sr isotopic composition of JA-2 was also measured at Okayama University, showing the 87Sr/ ⁸⁶Sr ratios of 0.706295 \pm 0.000024 (2SD, n = 9) (**Table 2**). These results indicate that our technique was robust enough to measure 87Sr/86Sr ratios.

TABLE 2 | Sr isotopic compositions of JA-2 and JB-1b.

Sample	Instrumentation and Reference	n	⁸⁷ Sı	′/ ⁸⁶ Sr
			Mean	2SD
JA-2	MC-ICP-MS (NIES)	7	0.706315	0.000022
	TIMS (Kumamoto Univ.)	1	0.706311	0.000007 ^a
	TIMS (Kumamoto Univ.)	1	0.706325	0.000007 ^a
	TIMS (Okayama Univ.)	9	0.706295	0.000024
	Miyazaki and Shuto (1998) ^b	5	0.706331	0.000013
JB-1b	MC-ICP-MS (NIES)	8	0.704093	0.000023
	TIMS (Kumamoto Univ.)	1	0.704083	0.000007 ^a
	TIMS (Kumamoto Univ.)	1	0.704084	0.000006 ^a
	Yuhara et al. (2000) ^b	13	0.704095	0.000012

^aThe analytical error is described in 2SE. 150 ratios (15 cycles x 10 blocks) were taken in a single measurement.

^bReported ⁸⁷Sr/⁸⁶Sr ratios of the NIST SRM 987 are 0.710251 ± 0.000004 (2*a*_m, *n* = 51) in Miyazaki and Shuto (1998), and 0.710251 in Yuhara et al. (2000). The ⁸⁷Sr/⁸⁶Sr ratios of the samples are corrected for interlaboratory bias by adjusting the mean value of the NIST SRM 987 standard run with the samples to the value of 0.710248.

TABLE 3 | Sr isotopic ratios of NIES CRM No. 28 Urban Aerosols.

Instrumentation		Pretreatment		Bottle	Number of	Number of	⁸⁷ Sr/ ⁸⁶ Sr	
		Decomposition	Sr separation	No.	subsampling	measurements for each subsample	Mean	2SD
MC-ICP-MS	Nu Plasma II	Hotplate with HNO ₃ /	Sr spec resin	044	3	2	0.710221	0.000014
(NIES)		HCIO ₄ /HF mixture		375	3	2	0.710233	0.000024
				597	3	2	0.710228	0.000015
						Mean	0.710227	0.000019
TIMS	TRITON	Hotplate with HNO ₃ /	Sr spec resin	044 ^a		4	0.710229	0.000013
(Kumamoto		HCIO ₄ /HF mixture		375 ^a		4	0.710226	0.000012
Univ.)				597 ^a		4	0.710233	0.000009
						Mean	0.710229	0.000011
TIMS (Okayama Univ.)	Finnigan MAT 262	High-pressure bomb with HNO ₃ /HF/HClO ₄	Cation exchange resin, AG50 × 12, 200–400 mesh	035	1	12	0.710226	0.000019

^aSample aliquots, decomposed and pretreated at NIES, were analyzed at Kumamoto University. Average internal precisions were ± 0.000013, ± 0.000007, and ± 0.000009 (2SE) at NIES, Kumamoto University, and Okayama University, respectively.

Homogeneity of Sr Isotopic Compositions for NIES CRM No. 28

Table 3 shows the Sr isotopic ratio of NIES CRM No. 28. Withinand between-bottle were evaluated using MC-ICP-MS at NIES, yielding ⁸⁷Sr/⁸⁶Sr isotopic ratios of 0.710227 \pm 0.000019 (2SD, n = 18). The uncertainty of the Sr isotopic values is an expanded uncertainty determined using a coverage factor k = 2, which corresponded to a confidence interval of ~95%. To investigate the homogeneity of the isotopic results in CRM (**Table 4**), the ⁸⁷Sr/⁸⁶Sr isotopic ratios were tested using a one-way analysis of variance (ANOVA). The between-bottle variation was not statistically significant (p > 0.05 and $F_{calculated value} < F_{critical value}$) as evaluated by one-way ANOVA. Therefore, when applied to the Sr isotopic ratios presented in this study, the CRM was homogeneous.

Interlaboratory Studies

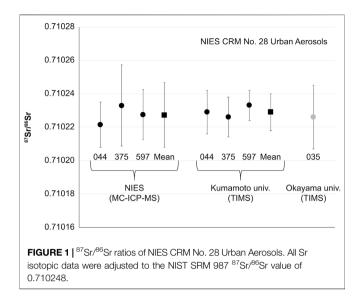
For the MC-ICP-MS vs. TIMS comparison, the Sr isotopic ratios of the three CRM samples (bottle No. 044, 375, and 597), which had been digested and processed for Sr separation at NIES, were measured by TIMS at the Kumamoto University to confirm the consistency of the

TABLE 4	ANOVA dat	a from the	homogeneity	study for	the Sr isotope.
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	F value	p value	F critical value	s _{bb} (%)	u _{bb} (%)	
⁸⁷ Sr/ ⁸⁶ Sr	2.403	0.1244	3.682	0.0006	0.0002	

 87 Sr/ 86 Sr isotopic ratios due to different analytical instruments: NIES was 0.710227 ± 0.000019 (2SD, n = 18) and Kumamoto University was 0.710229 ± 0.000011 (2SD, n = 12) (Figure 1; Table 3).

For sample decomposition and Sr separation, different methods were performed to evaluate bias during sample digestion and Sr separation at NIES and Okayama University. During the decomposition of the HNO₃/HClO₄/HF mixture using a hotplate at NIES, undissolved residues were observed, whereas a high-pressure bomb with HNO₃/HF/HClO₄ at Okayama University achieved complete digestion. The ⁸⁷Sr/⁸⁶Sr isotopic ratio obtained at Okayama University was 0.710226 \pm 0.000019 (2SD, n = 12) (**Figure 1; Table 3**). Despite the incomplete dissolution at NIES, the ⁸⁷Sr/⁸⁶Sr



isotopic ratio was consistent with the latter. Thus, the NIES CRM No. 28 Urban Aerosols was homogeneous enough for the Sr isotopic measurement of 100–200 mg subsamples using the methods described in *Sample Preparation*. This CRM will be of great value for the analytical quality assurance of environmental monitoring studies of PM.

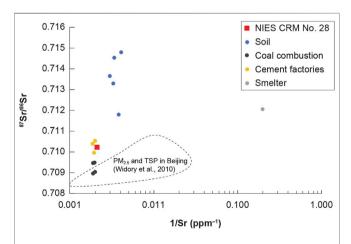
Potential CRM Emission Sources

The present study investigates the use of ⁸⁷Sr/⁸⁶Sr isotope systematics to help determine the origin of atmospheric aerosols. By characterizing the isotopes of ambient PM, potential sources of pollution near sampling sites can be identified, and their contribution to the contents of PM can be estimated. As a test, we compared the ⁸⁷Sr/⁸⁶Sr ratio of NIES CRM No. 28 with those determined by a previous source determination study. Widory et al. (2010) analyzed 63 samples of ambient PM_{2.5} and 23 samples of ambient total suspended particle (TSP) collected at various locations around Beijing from September 2005 to September 2006. The ⁸⁷Sr/⁸⁶Sr ratios of PM₂₅ and TSP ranged from 0.7085 to 0.7108, with the TSP having a larger variation. They also reported the ⁸⁷Sr/⁸⁶Sr ratios of potential emission sources near Beijing; coal combustion yielded aerosols with the lowest ⁸⁷Sr/⁸⁶Sr (0.708970-0.709492), smelter-derived particles produced the greatest radiogenic value (0.712064), and cement factories created particles with intermediate ⁸⁷Sr/⁸⁶Sr (0.709963-0.710528). They concluded that Sr in atmospheric PM in Beijing was mainly controlled by coal combustion and to a lesser extent by cement plants and/or smelters. Although sampling periods of the CRM, PM2.5 and TSP overlapped only for 1 year, and the 87 Sr/86 Sr ratio of CRM was a 10-years integral, the CRM ⁸⁷Sr/⁸⁶Sr ratio plotted within the range of the reported ratios for PM2.5 and TSP and near those of cement factories and coal combustion (Figure 2). The ⁸⁷Sr/⁸⁶Sr ratio of potential end-members, soil (0.711784-0.714797) and smelters, were significantly lower than that of the CRM.

To gain more insight into the sources of the CRM, enrichment factors (EFs) were calculated related to the Earth's upper

continental crust (Taylor and McLennan, 1995) with Fe as the reference element (3.50%). The mass fractions of the CRM metallic elements were reported by Mori et al. (2008). EFs >10 of the selected element (regarding anthropogenic sources), were higher in the order of Sb (120.5), As (72.1), Cd (68.5), Pb (24.2), Mo (22.7), and Zn (19.3) (**Table 5**). Some previous studies have reported that the major source of Sb and Cu in urban atmospheres was brake abrasion particles (Hjortenkrans et al., 2007; Iijima et al., 2007). However, CRM Cu EFs were not as high as Sb ($EF_{Cu} = 5.0$). High Sb EFs were also found in aerosols collected in Beijing from 2001 to 2006, and coal combustion was suggested as a possible extra source of Sb (Okuda et al., 2008). In addition, coal combustion has been considered a major source of As, the second-highest CRM EF in aerosols (Kowalczyk et al., 1798; Wang et al., 1999).

As a preliminary result, another fingerprint tracer, Pb isotopic composition, was obtained for the CRM at the Institut des Sciences Analytiques et de Physico-chimie pour l'Environnement et les Matériaux using MC-ICP-MS (Nu Instrument, UK) combined with a desolvator nebulizer unit (DSN-100, Nu Instrument). The Pb isotopic compositions of the NIES CRM No. 28 and other aerosol samples are summarized in Table 6. Combining the isotope systematics (206Pb/204Pb and 87Sr/86Sr) yield constraints for the source of the Beijing aerosols (see Figure 3), the results showed the impact of emissions from coal combustion and cement plants. The volume size distribution of the CRMs showed that particles with a diameter of \sim 7 µm were present with the highest frequency (Mori et al., 2008). According to Pb concentrations and Pb isotopic compositions, Widory et al. (2010) suggested that PM_{2.5} samples are expected to be primarily influenced by activities such as lead refining, while the coarser TSP fraction is attributed to activities such as coal combustion or emissions from cement plants. The average 207Pb/206Pb value and Pb concentration of aerosols from lead refining plants, coal combustion, and cement factories are 0.8668 ± 0.031 and 920,967 \pm 165,969 ppm (2SD, n = 6), 0.8583 \pm 0.179 and 109 \pm



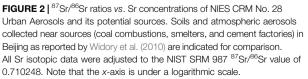


TABLE 5 | Enrichment factors (EFs) of >10 NIES CRM No. 28 and possible anthropogenic sources.

Element	EF	Possible anthropogenic sources	References			
Sb	120.5	Abrasion of vehicle brake linings, coal combustion, mining and smelting activities, electronic devices, road traffic, waste incineration, and the incineration of sewage sludge	Klumpp et al. (2009); Bech et al. (2012); Nriagu and Pacyna (1988); Okuda et al. (2008); Tian et al. (2012); Wang et al. (2003)			
As	72.1	Coal combustion, copper metallurgy, power plants, building materials, and electronics industries	Kowalczyk et al. (1798); Wang et al. (1999); Nriagu and Pacyna (1988); Christodouilidou et al. (2012); Cucu-Man and Steinnes (2013); Yang et al. (2015)			
Cd	68.5	Abrasion of tire treads and brake linings	Kummer et al. (2009)			
Pb	24.2	Gasoline, automobile emissions, abrasion of tire treads, brake linings, mining, Pb ore smelting, fertilizers, pesticides, and pigments	Yu et al. (2007); Kummer et al. (2009); Ribeiro de Souza et al. (2012)			
Мо	22.7	Smelting, chemical industries, electronics industries, mining, pharmaceuticals, and pesticides	Shan et al. (2013); Brankov et al. (2012)			
Zn	19.3	Vehicle components (traffic exhaust, tire and brake wear, and lubricating motor oil), fossil fuel combustion, electroplating, building materials, and electronics industries	Fujiwara et al. (2011); Huston et al. (2012); Mendiguchia et al. (2007); Robert-Sainte et al. (2009)			

EFs were calculated relative to the Earth's upper continental crust (Taylor and McLennan, 1995), and Fe was used as the reference element in this study (3.50%).

TABLE 6 | Pb isotopic compositions for aerosols.

Reference	²⁰⁶ Pb/ ²⁰⁴ Pb	²⁰⁷ Pb/ ²⁰⁴ Pb	²⁰⁸ Pb/ ²⁰⁴ Pb	²⁰⁶ Pb/ ²⁰⁷ Pb	²⁰⁸ Pb/ ²⁰⁷ Pb
This study	17.859	15.567	38.000	1.14719	2.44105
(2SD, n = 3)	0.008	0.005	0.017	0.00023	0.00026
Bing-Quan et al. (2002)	17.78	15.49	37.85	1.148	2.444
Chen et al. (2005)				1.11	2.434
Chen et al. (2005)				1.147	2.435
Widory et al. (2010)	18.064			1.161	
Widory et al. (2010)	18.879			1.155	
	18.873			1.154	
	This study (2SD, n = 3) Bing-Quan et al. (2002) Chen et al. (2005) Chen et al. (2005) Widory et al. (2010)	This study 17.859 (2SD, n = 3) 0.008 Bing-Quan et al. (2002) 17.78 Chen et al. (2005) Chen et al. (2005) Widory et al. (2010) 18.064 Widory et al. (2010) 18.879	This study 17.859 15.567 (2SD, n = 3) 0.008 0.005 Bing-Quan et al. (2002) 17.78 15.49 Chen et al. (2005) Chen et al. (2005) Widory et al. (2010) 18.064 Widory et al. (2010) 18.879	This study 17.859 15.567 38.000 (2SD, n = 3) 0.008 0.005 0.017 Bing-Quan et al. (2002) 17.78 15.49 37.85 Chen et al. (2005) Chen et al. (2005) Vidory et al. (2010) 18.064 Widory et al. (2010) 18.879 15.49 15.49	This study 17.859 15.567 38.000 1.14719 (2SD, n = 3) 0.008 0.005 0.017 0.00023 Bing-Quan et al. (2002) 17.78 15.49 37.85 1.148 Chen et al. (2005) 11.11 1.147 1.147 Widory et al. (2010) 18.064 1.161 1.155

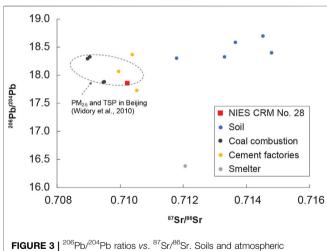


FIGURE 3 ²⁰⁰Pb/²⁰⁴Pb ratios vs. ⁶⁷Sr/⁸⁰Sr. Soils and atmospheric aerosols collected near sources (coal combustions, smelters, and cement factories) in Beijing as reported by Widory et al. (2010) are indicated for comparison. All Sr isotopic data were adjusted to the NIST SRM 987 ⁸⁷Sr/⁸⁶Sr value of 0.710248.

14 ppm (2SD, n = 4), and 0.8616 ± 0.278 and 21 ± 16 ppm (2SD, n = 3), respectively (Widory et al., 2010). The Pb concentration of lead refining plants is significantly high compared to that of CRM. The lead refining plants' derived aerosols cannot be the major

source, but a minor contribution to the CRM might be possible. A similarity in the Pb isotopic compositions of the CRM and TSPs in leaded gasoline vehicle exhaust from Shanghai was also reported (Chen et al., 2005). Since the complete ban on the use of alkyllead in 2000, atmospheric Pb emissions have significantly decreased. However, unleaded gasoline still contains a small amount of Pb, inherited from the crude oil, thus it could be a source of contamination (Wang et al., 2003; Bi et al., 2017). The original material for the CRM was recovered before this prohibition and during the phase-out of leaded gasoline, so the Pb isotopic ratio may record such environmental conditions.

CONCLUSION

NIES CRM No. 28 Urban Aerosols was originally prepared to certify mass fractions of major and minor elements. In this study, the Sr isotopic composition of the CRM was determined to provide an appropriate quality assurance/quality control tool for Sr isotopic analyses of atmospheric particles. To validate and ensure the accuracy of our method, secondary reference materials, JA-2 and JB-1b, were pretreated in the same way as the CRM, and Sr isotopic compositions were measured using MC-ICP-MS. According to our results regarding within- and betweenbottle variations of CRM subsamples, the CRM was sufficiently homogenous to be used for Sr isotopic measurements. As part of

an interlaboratory CRM study, same sample aliquots were measured using TIMS. The results confirmed the consistency of the isotopic ratio using two instruments (MC-ICP-MS *vs.* TIMS). We also confirmed the consistency of the Sr isotopic composition using different digestion methods (hotplates *vs.* digestion bombs) and Sr separation (Sr spec resin *vs.* cation exchange resin). The results of our isotopic analysis contribute to the quality assurance of environmental aerosol monitoring studies.

DATA AVAILABILITY STATEMENT

The original contributions presented in the study are included in the article, further inquiries can be directed to the corresponding author.

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AUTHOR CONTRIBUTIONS

AY, KY, DA, SB, and OD designed the research. AY, KN, MU, KO, KY, TS, KT, TK, and SB performed the analytical work. AY, KY, and TK wrote the manuscript. All the authors discussed the data and revised and approved the final form of the manuscript.

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